

**VISIBLE AND NEAR-IR REFLECTANCE SPECTRA FOR SMECTITE, SULFATE AND PERCHLORATE UNDER DRY CONDITIONS FOR INTERPRETATION OF MARTIAN SURFACE MINERALOGY.** R. V. Morris<sup>1</sup>, D. C. Golden<sup>2</sup>, D. W. Ming<sup>1</sup>, T. G. Graff<sup>2</sup>, R. E. Arvidson<sup>3</sup>, S. M. Wiseman<sup>3</sup>, K. A. Lichtenberg<sup>3</sup>, and S. Cull<sup>3</sup>. <sup>1</sup>ARES NASA Johnson Space Center, Houston TX 77058 ([richard.v.morris@nasa.gov](mailto:richard.v.morris@nasa.gov)), <sup>2</sup>ESCG-JW23, Houston, TX, <sup>3</sup>Washington University, St. Louis, MO.

**Introduction:** Visible and near-IR (VNIR) spectral data for the martian surface obtained from orbit by the MRO-CRISM and OMEGA instruments are interpreted as having spectral signatures of H<sub>2</sub>O/OH-bearing phases, including smectites and other phyllosilicates, sulfates, and high-SiO<sub>2</sub> phases [e.g., 1-4]. Interpretations of martian spectral signatures are based on and constrained by spectra that are obtained in the laboratory on samples with known mineralogical compositions and other physicochemical characteristics under, as appropriate, Mars-like environmental conditions (e.g., temperature, pressure, and humidity). With respect to environmental conditions, differences in the absolute concentration of atmospheric H<sub>2</sub>O can effect the hydration state and therefore the spectra signatures of smectite phyllosilicates (solvation H<sub>2</sub>O) and certain sulfates (hydration H<sub>2</sub>O) [e.g., 5-7].

We report VNIR spectral data acquired under humid (laboratory air) and dry (dry N<sub>2</sub> gas) environments for two natural smectites (nontronite API-33A and saponite SapCa-1) to characterize the effect of solvation H<sub>2</sub>O on spectral properties. We also report spectral data for the thermal dehydration products of (1) melanterite (FeSO<sub>4</sub>.7H<sub>2</sub>O) in both air and dry N<sub>2</sub> gas and (2) Mg-perchlorate (Mg(ClO<sub>4</sub>)<sub>2</sub>.6H<sub>2</sub>O) in dry N<sub>2</sub> environments. Spectral measurements for samples dehydrated in dry N<sub>2</sub> were made without exposing them to humid laboratory air.

**Methods:** VNIR reflectivity spectra (0.35-2.50 μm) were acquired with Analytical Spectral Devices (ASD) fiberoptic spectrometers configured with Mug lights. A FieldSpec3 Hi-Res spectrometer was used for measurements in laboratory air at ~25°C. A FieldSpec3 spectrometer, a hot plate (ambient to 400°C), and a dewpoint meter (Vaisala DRYCAP DM70) were co-located in a one-atmosphere glove box that can be kept under continuous purge using dry N<sub>2</sub> gas (from liquid N<sub>2</sub>) to provide a dry atmosphere. Spectral measurements on samples in the glove box were made in air (8000-12000 ppm H<sub>2</sub>O by volume) prior to starting the N<sub>2</sub> purge, after purging 24-650 hr at ~25°C (150-400 ppm H<sub>2</sub>O or 0.5-2.0% relative humidity), and, depending on the sample, after heating to 50-330°C on the hot plate and cooling to ~25°C. Because H<sub>2</sub>O evolves during heating, the time at temperature was governed by the time required for the H<sub>2</sub>O content of the gas at-

mosphere to return to <400 ppm.

Powder X-ray diffraction spectra (PANalytical X'Pert PRO) were obtained under ambient conditions for all samples within 5-10 min and ~1 hr after their removal from the glove box or after thermal treatment in air to determine mineralogical composition and phase stability.

**Results and Discussion:**

*Nontronite and Saponite.* Dehydration of nontronite and saponite under hyper-arid conditions results in a dramatic decrease in intensity of spectral features that require the presence of only the H<sub>2</sub>O molecule (e.g., at ~1.9 μm) relative to those that require the presence of only M-OH (e.g., at 2.29 and 2.31 μm for nontronite (M = Fe) and saponite (M = Mg), respectively (Fig. 1)). All M-OH spectral features between 2.0 and 2.5 μm are also better resolved with increasing loss of solvation H<sub>2</sub>O. Spectral features that are composites of H<sub>2</sub>O-only and M-OH-only absorptions will also decrease in intensity and can shift in position if the individual peaks occur at different wavelengths. Saponite is an example, where a broad band centered near 1.41 μm (Fig. 1b, blue) under humid conditions is sharp and centered at 1.39 μm for measurements under hyper-arid conditions (red).

The enhancement of M-OH relative to H<sub>2</sub>O spectral features in Fig. 1 is in agreement with previous studies of smectites [e.g., 6-8]. The magnitude of the enhancements cannot be directly compared, however, because environmental H<sub>2</sub>O concentrations are not reported in all the earlier work.

If climatic conditions on Mars favor loss of solvation H<sub>2</sub>O over geologic time, these results imply that spectral features for M-OH should dominate over those for H<sub>2</sub>O for smectites and other phases that can exchange H<sub>2</sub>O with the martian atmosphere. We suggest that dehydration of solvation H<sub>2</sub>O from smectites by heating at ~110°C for laboratory timescales is a reasonable surrogate for dehydration of solvation H<sub>2</sub>O over geologic timescales at lower temperatures.

*Fe-Sulfates.* The first thermal decomposition product of melanterite in both air and dry N<sub>2</sub> atmospheres is szomolnokite (FeSO<sub>4</sub>.H<sub>2</sub>O) (Fig. 2). The spectra are similar to literature szomolnokite spectra [e.g., 9], except that no spectral feature at ~0.43 μm from Fe<sup>3+</sup> is present. In humid air at higher temperatures (~240°C),

$\text{Fe}^{2+}$  oxidation occurs resulting in formation of a ferric sulfate hydroxide which is possibly hydrated (from XRD data). The spectral features from  $\text{Fe}^{3+}$ -OH are located at 1.49, 1.82, 2.23, and 2.37  $\mu\text{m}$ . This phase may occur naturally on Mars [10]. Under hyper arid (and presumably low  $\text{O}_2$ ) conditions, szomolnokite thermally altered to crystalline  $\text{FeSO}_4$  (from XRD) (Fig. 2b). The broad bands centered near 0.99 and 1.53  $\mu\text{m}$  are from  $\text{Fe}^{2+}$  electronic transitions.

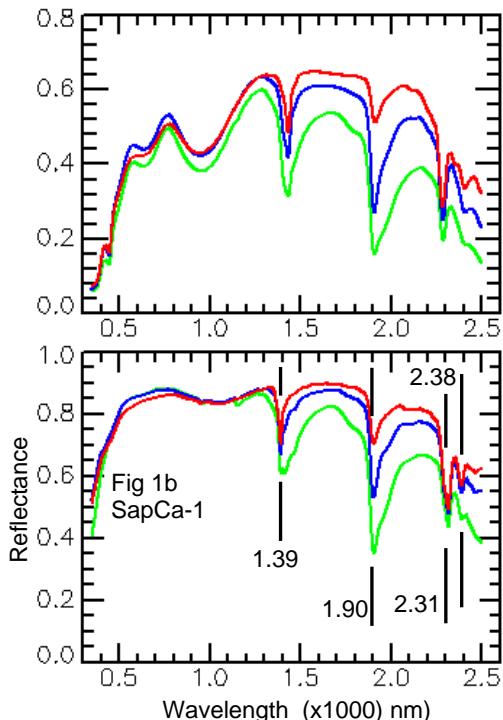


Fig. 1. (a) Nontronite (API-33A) spectra in air (green), after 24 hr  $\text{N}_2$  purge at 25°C with final  $\text{H}_2\text{O} = 266$  ppm (blue), and after 24 hr  $\text{N}_2$  purge at 110°C with final  $\text{H}_2\text{O} = 200$  ppm (red). (b). Saponite SapCa-1 spectra at ambient conditions (green), after 24 hr  $\text{N}_2$  purge at 25°C with final  $\text{H}_2\text{O} = 266$  ppm (blue), and after 24 hr  $\text{N}_2$  purge at 110°C with final  $\text{H}_2\text{O} = 200$  ppm (red).

*Mg-Perchlorate.* The three most intense spectral features for fully and partially hydrated Mg-perchlorate are located at 1.42, 1.91, and 2.38  $\mu\text{m}$  (Fig. 3). Because these positions approximate those for hydrous ferric sulfate, it is possible that locations previously interpreted as polyhydrated sulfates are actually mixed (and hydrated sulfate and perchlorate salts).

**References:** [1] Bibring *et al.* (2006). *Science*, 312, 400. [2] Poulet *et al.* (2005), *Nature*, 438, 623. [3] Mustard *et al.*, *Nature*, doi:10.1038/nature07097, 305. [4] Milliken *et al.* (2008) *Geology*, 36, 847. [5] Bishop *et al.* (1994) *Clays Clay*

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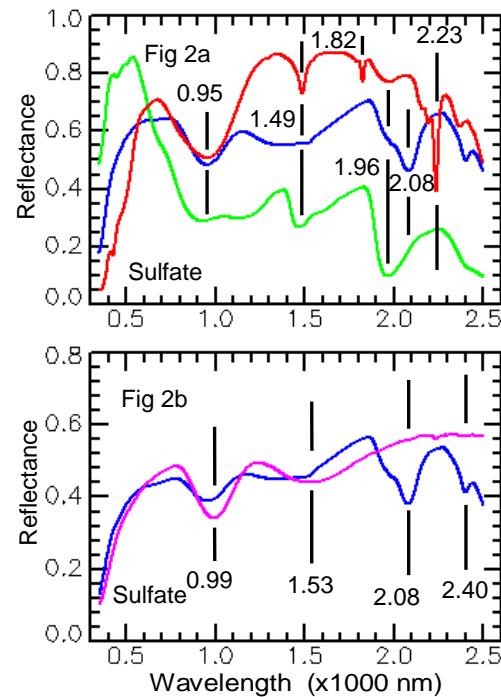


Fig. 2. (a) Spectra of melanterite ( $\text{FeSO}_4 \cdot 7\text{H}_2\text{O}$ ) (green) and szomolnokite ( $\text{FeSO}_4 \cdot \text{H}_2\text{O}$ ) (blue) and ferric sulfate hydroxide (red) obtained by thermal dehydration in air. (b) Spectra of szomolnokite (blue) and anhydrous  $\text{FeSO}_4$  (magenta) obtained as thermal dehydration products of melanterite under dry conditions at 210°C and 310°C, respectively.

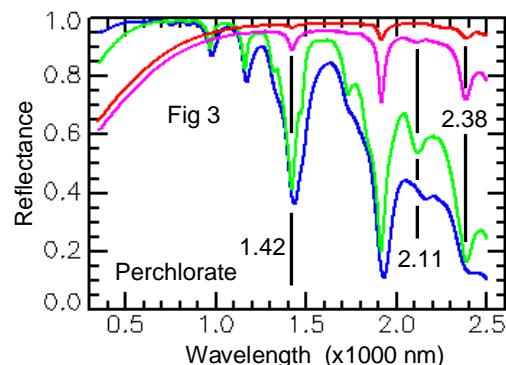


Fig. 3. Spectra of unheated  $\text{Mg}(\text{ClO}_4)_2 \cdot 6\text{H}_2\text{O}$  (blue) and its thermal dehydration products in dry  $\text{N}_2$  at 210°C (green) and 330°C (red and magenta).